

# Optical characterization of chromophoric dissolved organic matter at eutrophic and oligotrophic parts of a semi-enclosed bay (İzmir, Aegean Sea)

## Yarı-kapalı bir körfezin (İzmir, Ege Denizi) ötrofik ve oligotrofik kısımlarında kromoforik çözünmüş organik maddenin optik karakterizasyonu

Hakan Alyürük<sup>1\*</sup> • Aynur Konaş<sup>2</sup>

<sup>1</sup>Dokuz Eylül University, Institute of Marine Sciences and Technology, 35340, İnciraltı, İzmir, Turkey

<https://orcid.org/0000-0001-8632-4281>

<sup>2</sup>Dokuz Eylül University, Institute of Marine Sciences and Technology, 35340, İnciraltı, İzmir, Turkey

<https://orcid.org/0000-0002-6273-1568>

\*Corresponding author: [hakan.aluyuruk@deu.edu.tr](mailto:hakan.aluyuruk@deu.edu.tr)

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**Abstract:** Optical characterization of chromophoric dissolved organic matter (CDOM) from İzmir Bay (Aegean Sea) waters was investigated. For sampling, surface and subsurface seawater from 7 stations were collected in summer 2015. Excitation-emission matrix (EEM) spectra of each sample were recorded on a fluorescence spectrophotometer. The results showed that dissolved organic carbon (DOC) concentrations and EEM peaks were increased from the outer bay to inner bay stations. EEM peaks indicated the presence of both humic-like and protein-like components which were higher at middle-inner bays than outer bay. Spearman's rank correlation coefficients for EEM peak intensities and DOC concentrations were highly positive ( $p < 0.05$ ). HIX found between 0.73-3.51, whereas BIX ranged from 0.31 to 0.96 in the bay. Humification degree of CDOM in the middle-inner bays were higher compared to outer bay stations. High HIX values in the middle-inner bays could be linked to the presence of Melez stream (heavily polluted), other streams, rain run-offs and maritime activities at İzmir Bay. High BIX values in the middle-inner bays indicated presence of freshly produced DOM from bacterial origin. Optical characterization of CDOM could be used for tracing fluorescent DOM components and determining different DOM sources (autochthonous or allochthonous) in further studies.

**Keywords:** Chromophoric dissolved organic matter, dissolved organic carbon, fluorescence, seawater, İzmir Bay

**Öz:** Bu çalışmada, İzmir Körfezi (Ege Denizi)'nden alınan deniz suyunda kromoforik çözünmüş organik madde (KÇOM)'nin optik karakterizasyonu araştırılmıştır. Deniz suyu örnekleri 7 istasyonda yüzey ve yüzey-altı derinliklerden 2015 yılı yaz mevsiminde toplanmıştır. Her bir örneğin uyarma-emisyon matris (UEM) spektrumu bir floresans spektrofotometresi aracılığıyla kaydedilmiştir. Sonuçlar, çözünmüş organik karbon (ÇOK) ve UEM pik şiddetlerinin dış körfezden iç körfeze doğru arttığını göstermiştir. UEM pikleri, hem humik asit benzeri hem de protein benzeri organik madde bileşenlerinin orta-iç körfezlerde dış körfeze kıyasla daha yüksek olduğunu ortaya koymuştur. Spearman'ın sıralama korelasyonu testleri sonucunda UEM pik şiddetleri ve ÇOK konsantrasyonları arasında yüksek pozitif ilişki bulunduğu saptanmıştır ( $p < 0.05$ ). Körfez genelinde humikleşme indeksi (HI) 0.73-3.51 aralığında, biyolojik indeks (BI) ise 0.31-0.96 aralığında değişim göstermiştir. KÇOM'a ait humikleşme derecesi orta-iç körfezlerde dış körfeze kıyasla daha yüksek tespit edilmiştir. Orta-iç körfezlerde gözlenen yüksek HI değerlerinin yüksek kirlilik yükü taşıyan Melez çayının varlığı, iç körfeze ulaşan diğer yüzey sularının varlığı, yağmur ile denize sürüklenebilen karasal organik maddeler ve İzmir Körfezi'ndeki denizcilik aktiviteleri ile ilişkili olduğu düşünülmektedir. Orta-iç körfezlerde gözlenen yüksek BI değerleri ise özellikle bakteriyel kökenli ve yeni üretilmiş çözünmüş organik madde üretimi ile ilişkilendirilebilir. KÇOM'un optik karakterizasyonu, gelecek çalışmalarda çözünmüş organik madde içerisindeki floresans özellikteki bileşenlerin takip edilebilmesi ve çeşitli organik madde kaynaklarının (otokton veya allokton) belirlenmesi amacıyla kullanılabilir.

**Anahtar kelimeler:** Kromoforik çözünmüş organik madde, çözünmüş organik karbon, floresans, deniz suyu, İzmir Körfezi

## INTRODUCTION

Dissolved organic matter (DOM) is a heterogeneous compound pool that is composed of complex organic molecules within seawater. DOM could be originated from terrestrial processes (allochthonous) or *in situ* marine (autochthonous) sources (Hedges, 2002; Libes, 2009). DOM pool contains vast amounts of biomolecules at different chemical structures and molecular sizes. Humic acids, fulvic acids, lignins, amino acids, carbohydrates, lipids, fatty acids, and sterols are such examples to most abundant biomolecules

in DOM. Primary production and exudates of phytoplankton, sloppy feeding of metazoan grazers, viral cell lysis, egesta of protists and metazoans, and extracellular hydrolysis of POM by bacteria are the main sources of *in situ* DOM production (Aparicio et al., 2016; Brussaard, 2004; Motegi et al., 2009; Ortega-Retuerta et al., 2009; Ridgwell and Arndt, 2015; Romera-Castillo et al., 2011a; Sala and Güde, 2004; Sarmiento et al., 2013; Zeri et al., 2014). Following its release to the marine environment, DOM is constantly circulated, transferred,

and subjected to transformations within the marine environment. The factors controlling its transformations could be photochemical processes (Mopper et al., 1991; Santos et al., 2014; Sulzberger and Durisch-Kaiser, 2009; Vähätalo and Wetzel, 2004; Zhang et al., 2013), sorption in sediments by flocculation (Cauwet, 2002), sorption by sinking particles (Carlson and Hansell, 2015; Hansell et al., 2009), sorption by metal-oxides (Couturier et al., 2016), physical processes (Boyd and Osburn, 2004; Cauwet, 2002; Dixon et al., 2014), and bacterial processes (Boyd and Osburn, 2004; Nelson et al., 2004; Santos et al., 2014; Vähätalo and Wetzel, 2004). Significant contributions of anthropogenic inputs to the DOM pool in coastal waters were also reported (Hong et al., 2005; Sun et al., 2014; Tedetti et al., 2011; Tzortziou et al., 2015; Wang et al., 2014), which in turn, these may lead to eutrophication, hypoxia or harmful algal blooms (Anderson et al., 2002; Cloern, 2001; Conley et al., 2009; Davidson et al., 2012; Davidson et al., 2014; Heisler et al., 2008; Jessen et al., 2015; Korpinen and Bonsdorff 2015; Sellner et al., 2003). Therefore, understanding and monitoring of DOM sources and its transformations in coastal waters is of great importance for elucidating the fates of DOM components and the assessment of ecological status of the marine environment.

In recent studies, optical characterization of DOM has been used frequently as it provides information about autochthonous or allochthonous sources and its transformations by applying rapid and inexpensive methods (Coble et al., 1990; Kowalczuk et al., 2010; Kowalczuk et al., 2015; Lonborg et al., 2015; Lu et al., 2015; Nieto-Cid et al., 2006; Romera-Castillo et al., 2011b; Romera-Castillo et al., 2013; Yang et al., 2016; Zeri et al., 2014). The fraction of DOM that is able to absorb and/or emit light is called as chromophoric dissolved organic matter (CDOM). The light spectra of CDOM is recorded by applying simultaneous (or three dimensional) excitation-emission matrix (EEM) spectroscopy with a fluorometer. Parallel factor analysis (PARAFAC) is used to extract the characteristic peaks of humic-like or protein-like compounds (Murphy et al., 2013; Stedmon et al., 2003). Fluorometric characterization of CDOM, when combined with absorption measurements, not only provides information about its sources and transformation but also helps to understand its relations with bacterial and photochemical processes (Guo et al., 2007; Kowalczuk et al., 2003; Lu et al., 2015; Murphy et al., 2008; Yamashita and Tanoue, 2003).

Izmir Bay is located at the Eastern coast of the Aegean Sea. It has an L shaped structure, and its entrance is oriented to the north with its longer part. The hydrography of the Izmir Bay is influenced from several factors: exchanges between the atmosphere and the sea, exchange of water masses with the Aegean Sea, freshwater inputs with anthropogenic loads, topography of the bay, sea level changes, wind-driven circulations of water masses and winter convection (Sayin, 2003). Under the influence of these factors, water masses in the Izmir Bay could be divided into three different parts: Inner Bay water (anthropogenically polluted), Outer Bay water (the

water mass influenced from Gediz River and Aegean Sea, the upwelling water at Gülbahçe Bay, and the water mass located at salt production area), and Middle Bay water (connects Outer Bay to Inner Bay) (Sayin, 2003). Remarkable differences for DOC, Chl-a, dissolved inorganic nitrogen (DIN), and dissolved inorganic phosphorus (DIP) levels were reported at outer, middle and inner bay stations in the previous studies (Kontas et al., 2004; Kucuksezgin et al., 2005; Sunlu et al., 2012). Also, algal blooms and eutrophication have been observed in the inner bay that is under the influence of anthropogenic inputs (Ozkan et al., 2008; Sunlu et al., 2011). For example, temporarily, high abundances of *Ceratium furca* var. *eugrammum*, *Cylindrotheca closterium*, *Prorocentrum micans* and *Noctiluca scintillans* biomasses have been observed in the range of 2-5  $\mu\text{M C}$  (Sunlu et al., 2007). However, there are no studies on the optical characterization of DOM, and its possible sources in the Izmir Bay. The aim of this study was to investigate the optical characteristics of CDOM in middle-inner (eutrophic) and outer (oligotrophic) parts of Izmir Bay with dissolved organic carbon (DOC) levels, humification index (HIX) and biological index (BIX).

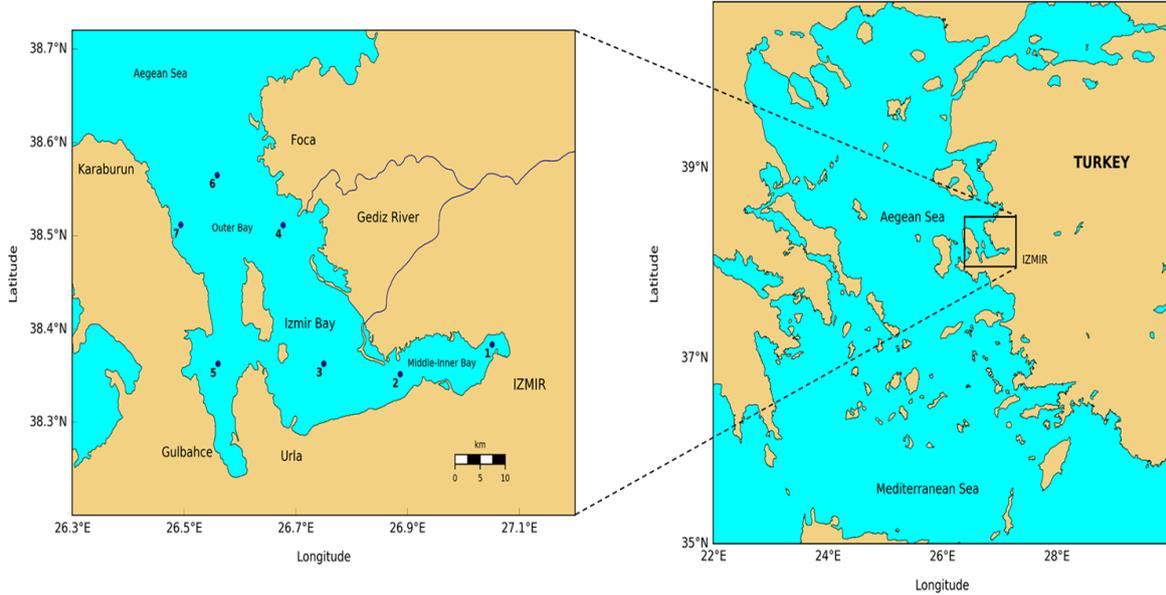
## MATERIAL AND METHODS

### Seawater sampling

The seawater samples were collected from surface and subsurface (5 m) depths at 7 stations in Izmir Bay (Figure 1) in summer 2015. The sampling stations were selected according to previous observations of physical and chemical characteristics of water masses in the bay (Kontas et al., 2004; Kucuksezgin et al., 2005; Sayin, 2003). The samples were collected with a 1.7 L Nansen bottle. The samples were immediately filtrated from 47 mm Whatman GF/F (0.7  $\mu\text{m}$ ) glass fiber filters and filtrated samples were stored frozen at  $-20\text{ }^{\circ}\text{C}$  until the analyses in the laboratory.

### Dissolved organic carbon analysis

The detection principle of DOC (measured as  $\text{CO}_2$ ) was based on the discoloration of buffered phenolphthalein solution proportionally to the  $\text{CO}_2$  concentration (Gershay et al., 1979; Schreurs, 1978). The analysis was performed on a continuous flow nutrient analyzer (San Plus, Skalar) according to instructions of the manufacturer (Cat.No: 311-412). First, 0.06 N sulfuric acid was added to the sample. Then, digestion reagent was added, and UV digestion was applied. Following the digestion, hydroxylammonium chloride solution was introduced and the  $\text{CO}_2$  was separated from reaction mixture with a gas dialysis membrane. The liberated  $\text{CO}_2$  was reacted with the 1% phenolphthalein buffer solution. Colorimetric reading was performed at 550 nm. Potassium hydrogen phthalate was used as organic carbon standard. Accuracy of the method was checked using potassium hydrogen phthalate at every 10 sample readings. Synthetic seawater including NaCl,  $\text{MgSO}_4$  and Milli-Q water was used as blank. The system was washed with Milli-Q water until the low and stable instrumental blank. The detection limit of the method was 16  $\mu\text{M C}$ .



**Figure 1.** Sampling stations at Izmir Bay, Aegean Sea

### Characterization of CDOM

EEM spectrums were recorded with an Agilent Cary Eclipse Fluorescence Spectrophotometer. EEM spectra were collected at excitation wavelengths of 230-500 nm and emission wavelengths of 250-600 nm with a spectral resolution of 2 nm. EEM spectra of Milli-Q water was subtracted from the EEM spectra of samples to remove scattering effects of water (Chari et al., 2012). Raman normalization was applied to normalize the data for comparability (Chari et al., 2012; Murphy et al., 2010; Stedmon et al., 2003) and fluorescence intensities were represented as Raman Units (RU). Raman normalization was performed by using Raman peak area (Eq. 1,  $\lambda_{ex}=350$  nm,  $\lambda_{em}=381-421$  nm) and dividing fluorescence intensity to Raman peak area (Eq. 2). EEM peaks of samples were determined according to Coble (1996) (Table 1). Post-processing of EEM spectrums were performed with modified PLOTEEM script in R (Lapworth and Kinniburgh, 2009; R Core Team, 2016). EEM peaks were extracted according to “algorithm-based approach” by selecting the peak points with maximum fluorescence intensities at defined emission regions (Korak et al., 2014).

$$A_R^{350} = \int_{381}^{426} I(\lambda_{em}) d\lambda_{em} \quad (1)$$

$$I(RU) = \frac{1}{A_R^{350}} I(AU) \quad (2)$$

where  $A_R^{350}$  indicates area under emission spectrum between 381 and 421 nm at excitation wavelength of 350 nm,  $I(\lambda_{em})$  represents fluorescence intensity as a function of emission wavelength,  $I(RU)$  represents fluorescence intensity of samples in RU,  $I(AU)$  represents raw fluorescence intensity of samples in Arbitrary Units (AU).

**Table 1.** Determination criteria of EEM peaks according to Coble (1996)

EEM peaks	Peak name (Wavelength range)
A	UV humic-like (Ex:260, Em:380-460)
B	Tyrosine-like, protein-like (Ex:275, Em:310)
C	Visible humic-like (Ex:350, Em:420-480)
M	Marine humic-like (Ex:312, Em:380-420)
T	Tryptophan-like, protein-like (Ex:275, Em:340)

### Calculations of fluorescence indexes

HIX and BIX are used to extract quantitative information on contributions of humic matter and autochthonous production to the fluorescence intensities of seawater samples. HIX was first introduced by Zsolnay et al. (1999) for estimating the humic matter content of DOM in soil samples. As a result of higher humification degree, C/H ratio and aromaticity was increased (Lüttig, 1986; Stevenson, 1982) and a shift to longer emission wavelengths was observed (Chen et al., 2011; Huguet et al., 2009; Tam and Sposito, 1993; Zsolnay et al., 1999). HIX is defined as the ratio of spectral area under emission wavelengths of 435-480 nm to emission wavelengths of 300-345 nm at excitation wavelength of 254 nm. HIX can be formulated as below (Eq. 3):

$$HIX = \frac{\int_{435}^{480} I(\lambda_{em}) d\lambda_{em}}{\int_{300}^{345} I(\lambda_{em}) d\lambda_{em}} \quad (3)$$

where  $I$  represents fluorescence intensity as a function of emission wavelength ( $\lambda_{em}$ ).

BIX was introduced by Huguet et al. (2009) for determination of the autochthonous biological activity and freshly produced DOM in natural water samples. While large BIX values indicate diagenetically unaltered DOM, numerator part of the fraction corresponds to microbially produced DOM, and denominator part represents highly decomposed DOM (Fellman et al., 2010; Lu et al., 2015; Wilson and Xenopoulos, 2009). It is calculated by dividing the fluorescence intensity at excitation wavelength of 310 nm and emission wavelength of 380 nm to the maximum emission intensity in the range of 420-480 nm at excitation wavelength of 310 nm (Eq. 4).

$$BIX = \frac{I_{\lambda_{ex}310/\lambda_{em}380}}{\max_{420 \leq \lambda_{em} \leq 480} I_{\lambda_{ex}310}} \quad (4)$$

where *I* represents fluorescence intensity.

### Statistical analyses

Statistical analyses were performed with R Statistical Computing Software, v3.3.1 (R Core Team, 2016). The Spearman's rank correlation test was performed between EEM peaks, HIX, BIX and DOC. Also, relationships between EEM peaks, HIX, BIX and DOC were investigated with linear regression analyses.

### RESULTS AND DISCUSSION

The results of DOC, EEM peak intensities of CDOM components and fluorescence indexes (HIX and BIX) were given in Table 2. DOC concentrations were found in the range of 35.3-244.2 µM throughout the bay. DOC levels in the middle-inner bays (stations 1 and 2) observed higher compared to outer bay (stations 3-7) (Figure 2a). DOC concentrations found in this study were similar to the previous reports (56.1-121 µM) for İzmir Bay (Kucuksezgin et al., 2005)

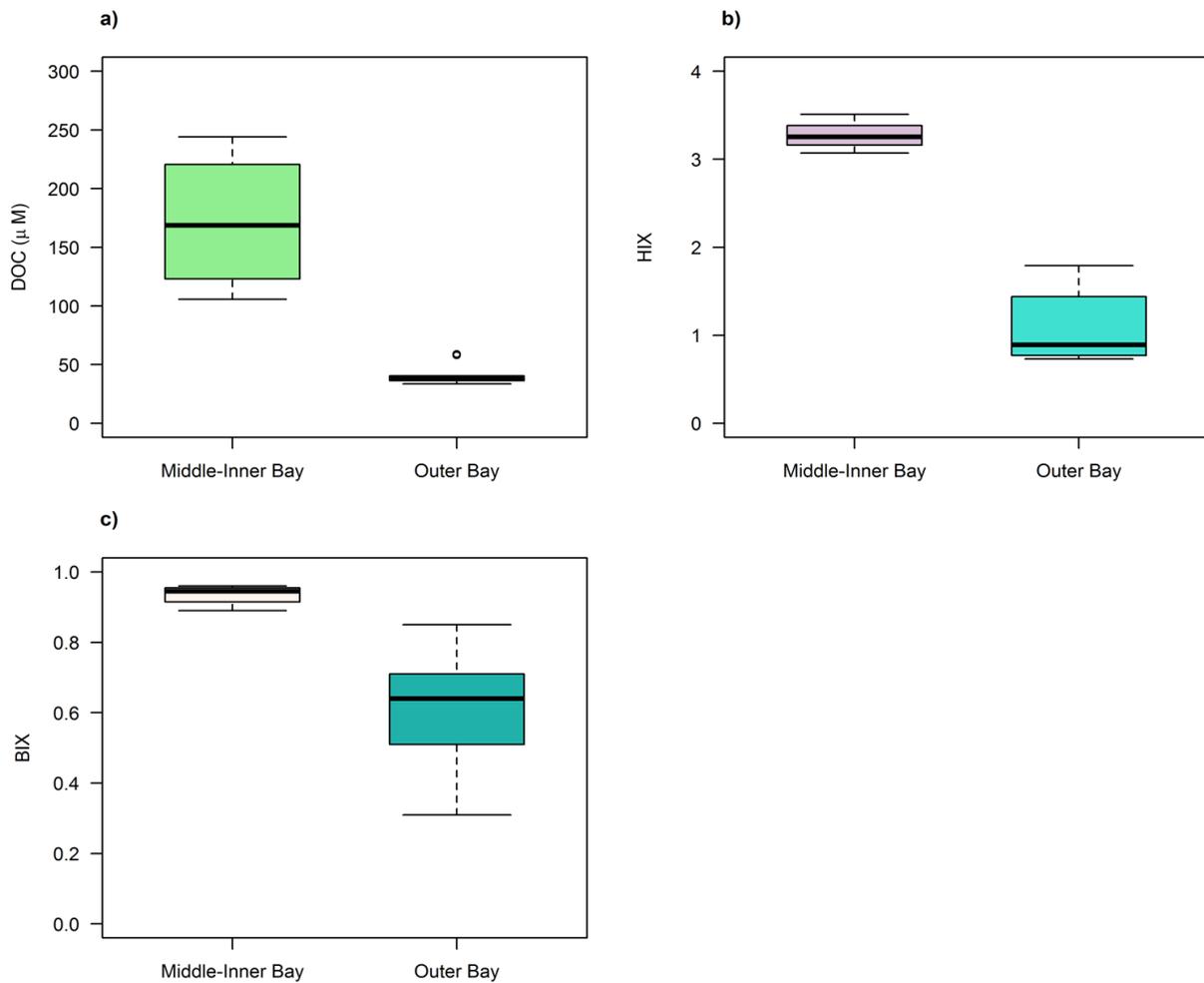


Figure 2. Distributions of DOC, HIX and BIX in the middle-inner and outer bays: a) DOC, b) HIX, and c) BIX

**Table 2.** DOC concentrations, fluorescent components of CDOM and descriptive fluorescence indexes (S: Surface, SS: Subsurface)

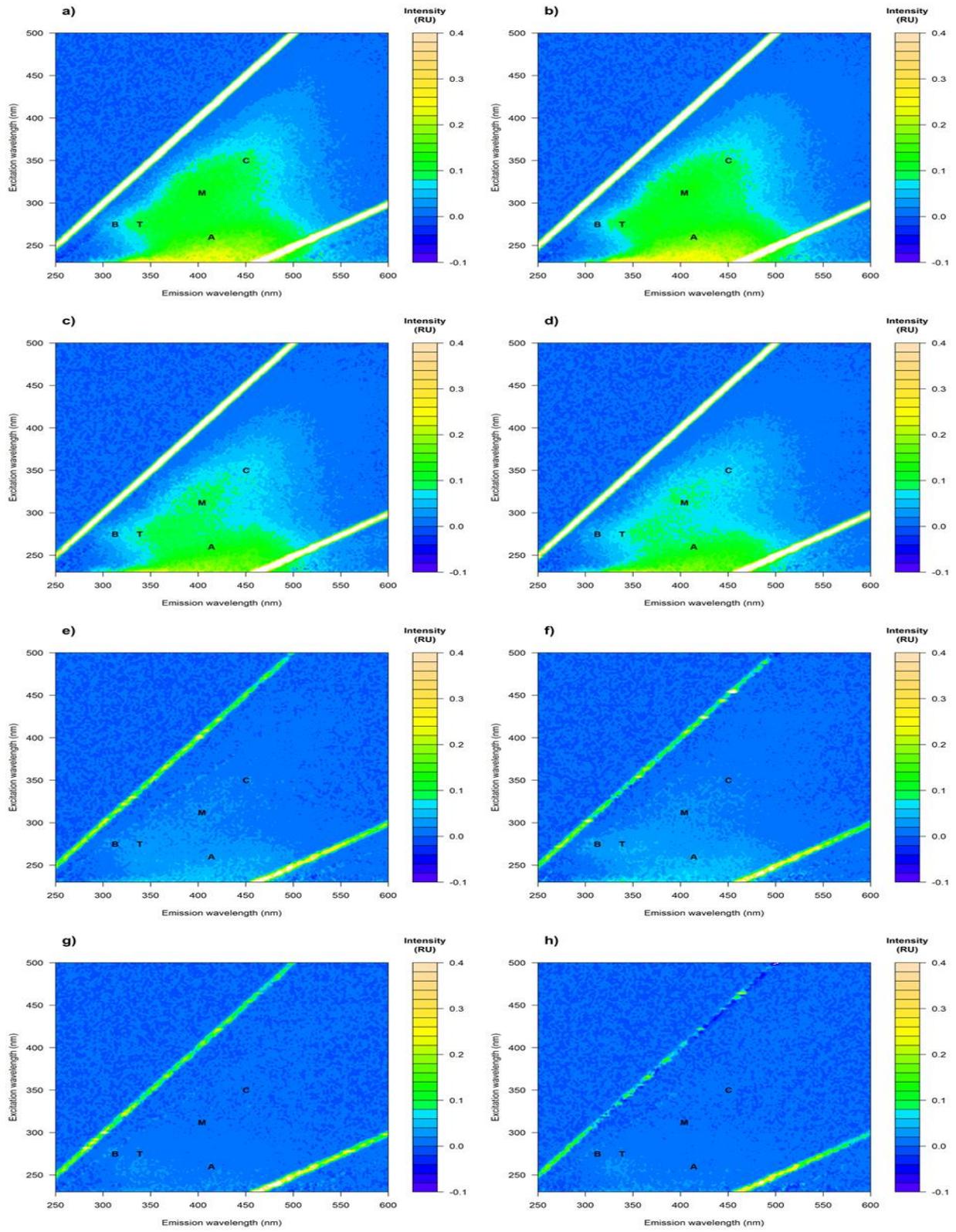
Station	DOC ( $\mu\text{M}$ )	Peak M	Peak A	Peak C	Peak T	Peak B	HIX	BIX
St1-S	244.2	0.13	0.18	0.11	0.11	0.05	3.51	0.94
St1-SS	206.6	0.14	0.18	0.10	0.09	0.06	3.26	0.95
St2-S	140.3	0.09	0.14	0.08	0.07	0.05	3.25	0.89
St2-SS	105.6	0.11	0.14	0.09	0.09	0.05	3.07	0.96
St3-S	58.9	0.03	0.05	0.04	0.04	0.02	1.79	0.85
St3-SS	42.7	0.03	0.05	0.02	0.02	0.02	1.58	0.68
St4-S	45.4	0.01	0.02	0.02	0.02	0.01	0.77	0.37
St4-SS	42.7	0.02	0.02	0.02	0.01	0.02	0.87	0.31
St5-S	38.6	0.03	0.04	0.02	0.03	0.02	1.08	0.60
St5-SS	46.7	0.03	0.03	0.02	0.02	0.01	1.44	0.73
St6-S	58.9	0.02	0.03	0.02	0.02	0.01	0.79	0.68
St6-SS	37.7	0.03	0.03	0.02	0.03	0.04	0.73	0.51
St7-S	38.6	0.02	0.03	0.02	0.01	0.01	0.77	0.57
St7-SS	35.3	0.01	0.02	0.02	0.01	0.01	0.91	0.71

Raman normalized fluorescence intensities of EEM peaks decreased from inner to outer bay stations (Table 2). EEM spectrums of the samples were given in Figures 3 and 4. Intensities of EEM peaks decreased in the order of A, M, C, T, and B at station 1 and 2. The intensities of peak A at station 3 were slightly higher compared to other outer bay stations. On the other hand, EEM peak intensities were very similar to each other at outer bay stations. Results indicated the presence of higher humic-like (peaks A, C, and M) and protein-like (peaks T and B) components at middle-inner bays than outer bay stations.

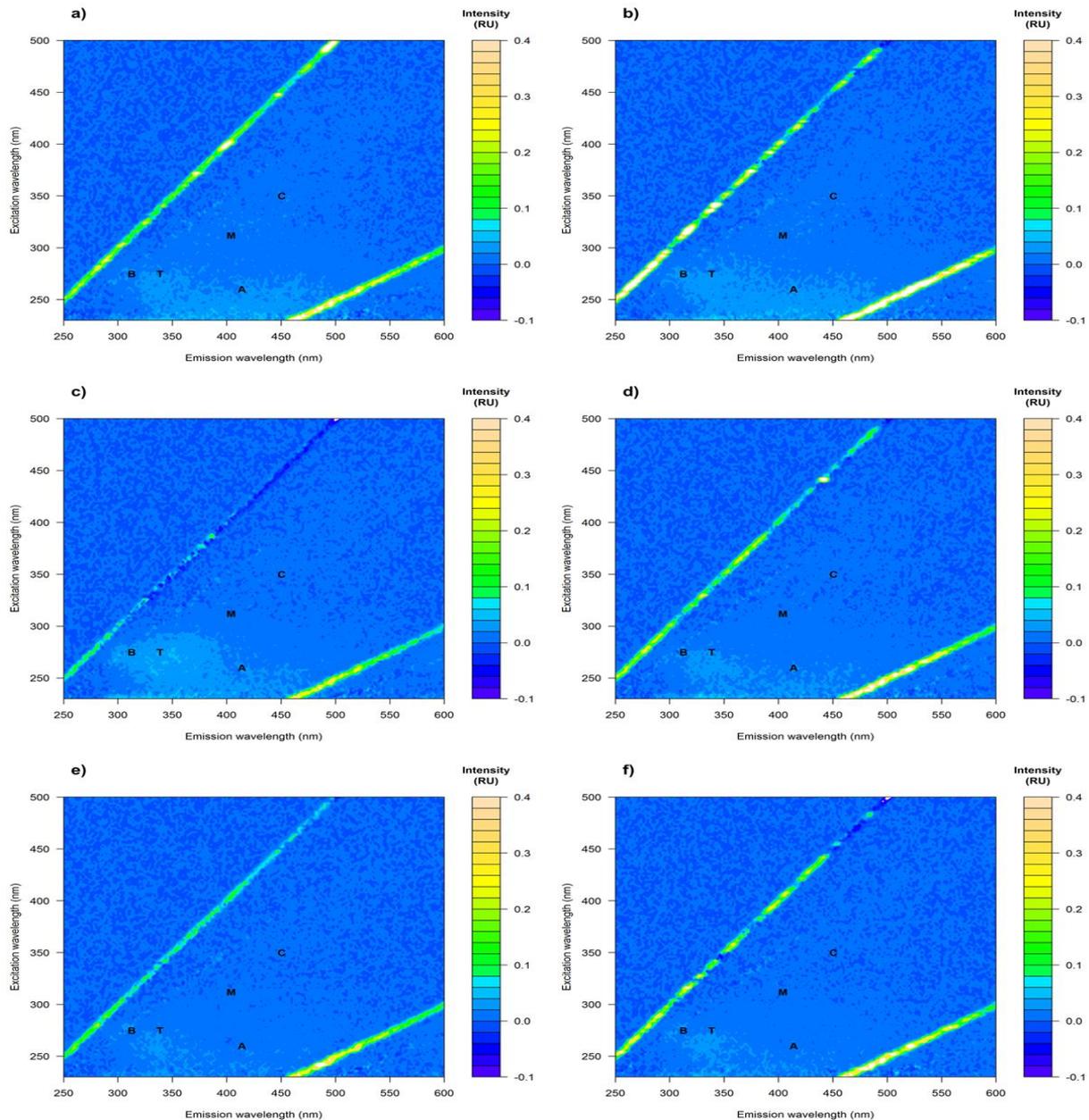
HIX and BIX indexes decreased from inner to outer bay stations (Figure 2b,c). HIX was between 0.73-3.51, whereas BIX ranged from 0.31 to 0.96 in the bay. According to scales of HIX defined by Huguet et al. (2009), HIX values in this study showed that the DOM in middle-inner bays (3.07-3.51) might be originated from biological or aquatic bacterial processes, but this character was weaker at outer bay stations (0.77-1.79).

According to BIX values (Huguet et al., 2009), DOM has strong autochthonous character at middle-inner bays (0.89-0.96) whereas DOM showed autochthonous character from low to strong levels at a wide range (0.31-0.85) in outer bay.

Observation of relatively high HIX values in the middle-inner bays could be linked to the presence of Melez stream (heavily polluted), other streams, rain run-offs and maritime activities at Izmir Bay. High HIX values at middle-inner bays might be related with humic matter inputs (humification degree) and stabilities of CDOM components (Bai et al., 2014; Chari et al., 2012; Huguet et al., 2009). Also, high BIX values indicated presence of freshly produced DOM from bacterial origin (Fellman et al., 2010; Lu et al., 2015; Wilson and Xenopoulos, 2009). According to HIX vs. BIX plot in Figure 5, it was possible to distinguish CDOM characteristics of middle-inner bay stations from outer bay. High HIX and BIX values for CDOM might be resulted from high anthropogenic inputs, physico-chemical characteristics and biological processes in middle-inner bays.



**Figure 3.** EEM plots for Station 1 to 4: a) St1-S, b) St1-SS, c) St2-S, d) St2-SS, e) St3-S, f) St3-SS, g) St4-S, and h) St4-SS (S: Surface, SS: Subsurface)



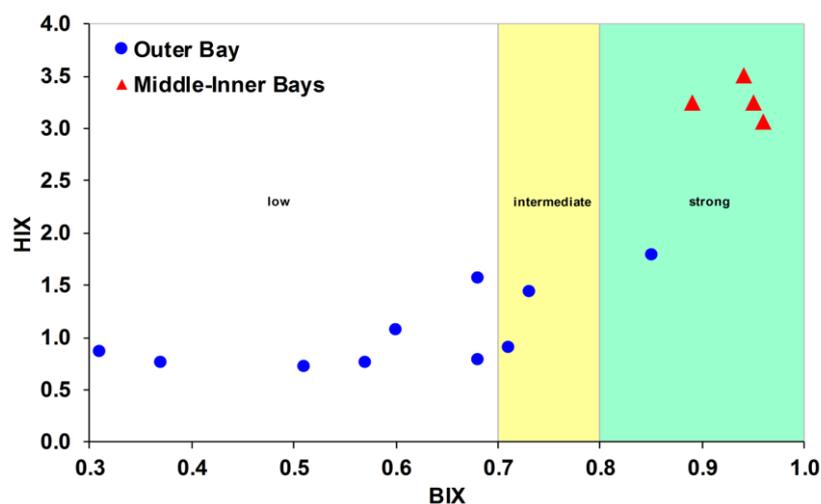
**Figure 4.** EEM plots for Station 5 to 7: a) St5-S, b) St5-SS, c) St6-S, d) St6-SS, e) St7-S, and f) St7-SS (S: Surface, SS: Subsurface)

Spearman's rank correlations between EEM peak intensities, HIX, BIX, total fluorescence intensities ( $\Sigma$ FL) and DOC concentrations were found highly positive ( $p < 0.05$ , Table 3). Linear relationships between EEM peaks, HIX, BIX and DOC concentrations at middle-inner and outer bays were given in Figure 6 and Table 3.

According to linear regression between DOC and  $\Sigma$ FL, the fluorescent fraction of DOM in the surface and subsurface waters of Izmir Bay were composed of highly fluorescent and slightly fluorescent fractions (Figure 6c). Also, humic-like (A, C, M) and tryptophan-like (T) EEM peaks were able to explain

more than 85% of variation in DOC concentrations. Linear relationships between HIX and DOC concentrations (explaining 80% variation in DOC) indicated that humic matter has an important contribution to DOM.

On the other hand, correlation ( $\rho = 0.719$ ,  $p = 0.004$ ) and linearity ( $r^2 = 0.484$ ,  $p = 0.005$ ) between BIX and DOC represented weaker contribution of autochthonous processes to DOM pool. As a result, DOM composition in Izmir Bay could be linked to humic matter inputs and *in situ* marine production processes.



**Figure 5.** Relationships between HIX and BIX. Low, intermediate and strong indicates autochthonous character of DOM (Huguet et al. 2009)

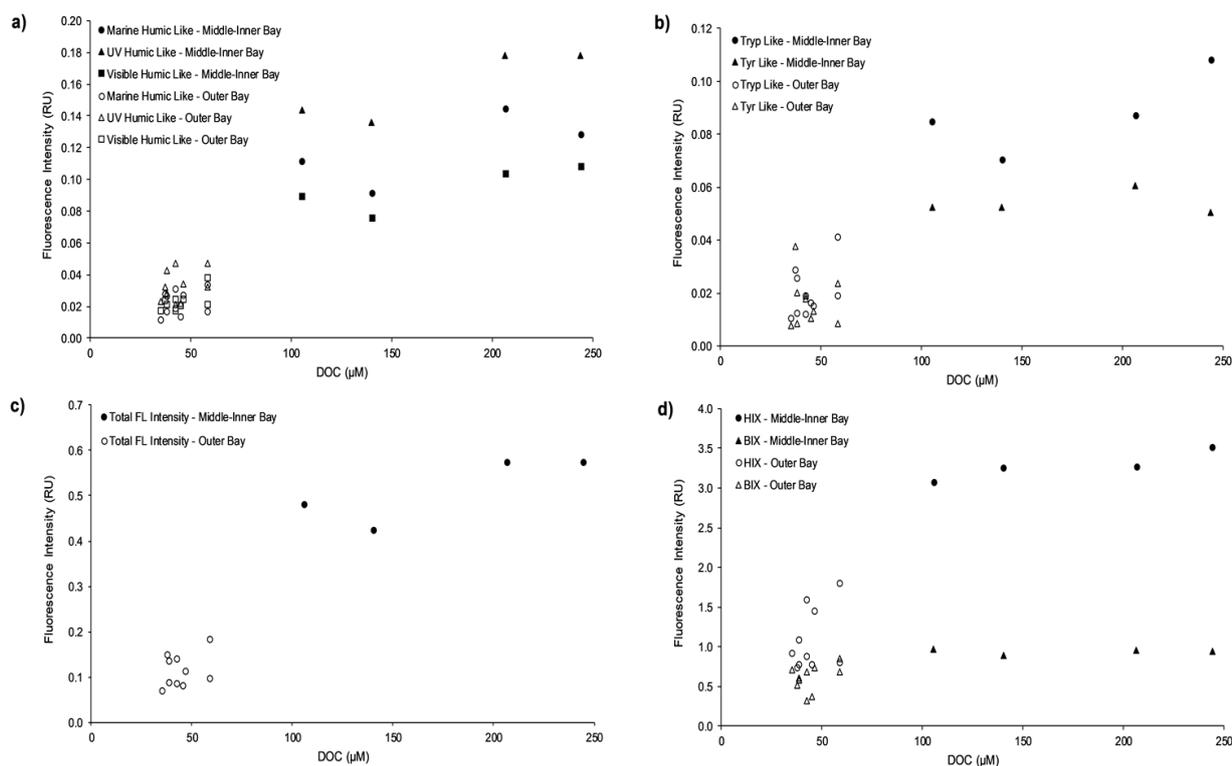
Spectral ranges of EEM peaks found in this study were very similar to the peaks of CDOM components previously reported in the literature (Penru et al., 2013; Retelletti Brogi et al., 2015; Su et al., 2015; Zeri et al., 2014). Excitation and emission wavelengths of peak regions and corresponding CDOM components were compared with other studies in Table 4. In the North Aegean Sea and Marmara Sea, Zeri et al. (2014) have identified three PARAFAC components within CDOM corresponding to UV and visible humic-like (C1), marine humic-like (C2) and tyrosine-like (C3) peaks. Similar humic-like and protein-like peaks were also present in studies at the Northwestern Mediterranean Sea (Penru et al., 2013) and Tyrrhenian Sea (Retelletti Brogi et al., 2015). Humic-like peaks have been reported at coastal environments under the influences of terrestrial inputs and low salinity waters (Coble, 1996; Stedmon and Markager, 2005). Marine humic-like peaks have been associated with increased phytoplankton activity

and autochthonous production (Coble, 1996; Murphy et al., 2008). Tyrosine and tryptophan-like peaks have been shown to originate from bacterial degradation of organic matter and these components are widely distributed along coastal waters (Coble, 1996; Stedmon et al., 2003; Yamashita et al., 2008).

Due to the low velocities of water masses inflowing and outflowing, inner part of İzmir Bay has been more heavily influenced from anthropogenic inputs. According to Sayin (2003), water mass renewal time for inner part was found as 3 months, whereas renewal time for outer bay was around 1-1.5 months. Similarly, Jiang et al. (2008) have reported the influence of water circulation on DOM transport. In middle-inner and outer bays, variations of EEM peak intensities, HIX and BIX (Figures 2-4) could be related to the influences of anthropogenic inputs and physico-chemical characteristics of water masses.

**Table 3.** Results of Spearman’s rank correlation tests and linear regression analyses between EEM peaks, Σ FL, HIX, BIX and DOC

	Correlations with DOC	Linear Regression Equations
Peak A	$\rho = 0.744, p = 0.002$	$[\text{DOC}] = 1044.9 [\text{A}] + 9.9, r^2 = 0.896, p = 0.000$
Peak B	$\rho = 0.575, p = 0.031$	$[\text{DOC}] = 2968.0 [\text{B}] + 1.0, r^2 = 0.654, p = 0.000$
Peak C	$\rho = 0.843, p = 0.000$	$[\text{DOC}] = 1826.9 [\text{C}] + 3.3, r^2 = 0.888, p = 0.000$
Peak M	$\rho = 0.701, p = 0.005$	$[\text{DOC}] = 1382.3 [\text{M}] + 12.5, r^2 = 0.873, p = 0.000$
Peak T	$\rho = 0.751, p = 0.002$	$[\text{DOC}] = 1828.4 [\text{T}] + 7.1, r^2 = 0.846, p = 0.000$
Σ FL	$\rho = 0.727, p = 0.003$	$[\text{DOC}] = 332.6 [\Sigma\text{FL}] + 5.8, r^2 = 0.877, p = 0.000$
HIX	$\rho = 0.774, p = 0.001$	$[\text{DOC}] = 56.1 [\text{HIX}] - 13.9, r^2 = 0.796, p = 0.000$
BIX	$\rho = 0.719, p = 0.004$	$[\text{DOC}] = 226.1 [\text{BIX}] - 75.9, r^2 = 0.484, p = 0.005$



**Figure 6.** Linear relationships between DOC and a) Marine Humic-like, UV Humic-like and Visible Humic-like peaks, b) Tryptophan-like and Tyrosine-like peaks, and c)  $\Sigma$ FL, and d) HIX and BIX

**Table 4.** Characteristics of the EEM peaks identified in this study and their comparison with those of other studies

EEM Peaks (Ex/Em) (This Study)	Marmara Sea; North Aegean Sea (Zeri et al., 2014)	Northwestern Mediterranean Sea (Penru et al., 2013)	Tyrrhenian Sea (Retelletti Brogi et al., 2015)	Southern Yellow Sea; East China Sea (Su et al., 2015)
A (260/380–460)	C1 (<260 (330)/464)	III (220–250/380–580)	P1 (250/400–500)	C2 (335/400)
B (275/310)	C3 (270/308)	I (220–250/280–332)	P4 (270/315)	-
C (350/420–480)	C1 (<260 (330)/464)	V (250–470/380–580)	P3 (350/450)	C1 (360/440)
M (312/380–420)	C2 (<250, 285/364)	IV (250–470/280–380)	P2 (315/419)	-
T (275/340)	-	II (220–250/332–380)	P5 (280/341)	C4 (280/360)

## CONCLUSION

Optical characterization of chromophoric dissolved organic matter in İzmir Bay was studied by excitation-emission matrix spectroscopy in this study. Fluorescence measurements indicated the presence of higher humic-like (peaks A, C, and M) and protein-like (peaks T and B) components at eutrophic middle-inner bays. DOC concentrations, EEM peak intensities and HIX/BIX values increased from outer to inner bay. High HIX and BIX values in the middle-inner bays could be explained by the presence of higher humification degrees and freshly produced DOM with bacterial origin. In middle-inner bays, CDOM composition might be influenced from nutrient

and DOM rich terrestrial inputs, and water renewal times. In conclusion, optical characterization of CDOM could be used for tracing fluorescent DOM components and determining different DOM sources (autochthonous or allochthonous) in further studies.

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